

EVROPSKÁ UNIE Evropské strukturální a investiční fondy Operační program Výzkum, vývoj a vzdělávání



Geochemistry on the Earth's surface for analytical geochemists

3b. Speed of processes

Tento učební materiál vznikl v rámci projektu Rozvoj doktorského studia chemie č. CZ.02.2.69/0.0/0.0/16_018/0002593

Questions

- 1. How can we determine which processes will take place spontaneously in nature?
- 2. What will tell us in which direction?
- 3. From what can we determine the equilibrium state of the system/process?
- 4. Which processes will not reach equilibrium, but will run (more or less) completely?

Reaction kinetics

- It studies the rate of reactions that are feasible from thermodynamic point of view.
- The rate of a reaction is basically the rate at which the reactants are consumed and the products are formed.
- How reactions occur is explained by the collision theory

Collision theory

- For the reaction to take place, several conditions must be met:
- 1. The reactant particles must collide.
- 2. The colliding particles must have sufficient energy to react.
- In some cases, particles need to collide with a specific orientation (e.g., a functional group).

Activation energy

- The minimum total kinetic energy that the particles must have for a reaction to occur.
- Each reaction has a different value of E_A .
- In the case of slow reactions, the activation energy is much higher than the average energy of the particles – collisions rarely lead to the reaction.
- Sometimes the activation energy can be supplied externally as an initiator, and then the reaction continues to be driven by the energy released during it.



Factors affecting reaction rate

- What factors will play a role in the final rate of chemical reactions?
- How will this be reflected in geological processes?

Factors affecting reaction rate

- There are several, but 4 generally apply to all reactions:
 - 1. Physical nature of reactants.
 - 2. Concentration.
 - 3. Temperature.
 - 4. The presence of catalysts.

Physical nature of reactants

- State, but also the size of the particles.
- Reaction rate: gas > liquid > solid
- Solids react on the surface:
 - The larger the surface, the faster the reaction.
 - The smaller the grains, the faster the reaction.
- Ex.: Consider the burning of the same amount of wood in different forms

Increase in surface area = increase in contact between reactants

Concentration

- An increase in reactants leads to an increase in the reaction rate.
- Substances burn faster in pure oxygen than in air.
- A higher concentration increases the chance of particle collision.
- The effect of concentration varies depending on the substance, there is no general relationship and each substance needs to be tested experimentally.
- E.g.: Influence of the content of CO₂, O₂ or organic acids in the soil on weathering and soilforming processes.

Temperature

- Increases the probability of collisions because the particles have higher kinetic energy.
- Particles with higher energy are closer to exceeding the activation energy (E_A).
- The reaction rate roughly doubles for every 10°C.
- **Ex.**: Food is refrigerated to slow down decomposition processes.
- Weathering in cold/warm climates.

The presence of catalysts

- A catalyst is a substance that increases the rate of a reaction but is itself unchanged after the reaction is complete.
- They provide alternative reaction pathways with lower activation energy.
- It is part of the reaction in the same way that a table is part of the lunch – it provides space and the right conditions, but the reaction does not consume it.



Reaction Progress

Homogeneous catalyst

• Same phase as the reactants, usually forming an intermediate.

A + B + cat -> A–B -> product + catalyst \swarrow cat

Heterogeneous catalyst

- A separated phase, usually a solid.
- It is assumed to provide the environment for the reaction.
 - The reactants bind to the surface of the solid phase at a certain angle and thus facilitate the reaction.

Use of catalysts

- Intensive use in the chemical industry in the production of chemicals -> low energy consumption during reactions is economically very advantageous.
- High temperatures are demanding and expensive!
- Exhaust gas cleaners in combustion engines of cars.
- Enzymes are important in biochemistry.
- Ozone depletion.

SPEED OF PROCESSES

A natural process

- We observe that the concentration of a substance in a solution changes over time
- There is a process going on
- How can I find out and express its speed?

Process speed

 We usually use the change in content over time – concentration, weight, amount of substance, volume

d[A]/d*t*

• The tangent of the concentration-time curve

Oth order reaction

A => B

d[A]/dt = -k $[A] = [A]_{o} - kt$

Oth order reaction



1st order reaction

A => B

d[A]/dt = -kA

 $[\mathbf{A}] = [\mathbf{A}]_{\mathbf{o}} * e^{-kt}$

1st order reaction



Adapted from Ryan (2014)

2nd order reaction

2A => BA + B = > C1/[A] = k * t + C $1/[A] - 1/[A_o] = k * t$

2nd order reaction



Dependence of k on temperature



Fig. 1.19 Effect of temperature on chemical weathering rate of rhyolite as a function of temperature. (Yokoyama and Banfield 2002. Reproduced with permission of Elsevier.)

Retention time

• Inverse of the rate constant

 $k = 1/\tau$

Example process

CALCITE DISSOLUTION

Dissolution/precipitation



$$\Omega = \frac{Q}{K} = \frac{[Ca^{2+}][CO_3^{2-}]}{K}$$

An empirical model

Dissolution $\Re_{d} = k_{d} (1 - \Omega_{cal})^{n_{d}}$ Precipitation $\Re_{p} = k_{p} (\Omega_{cal} - 1)^{n_{p}}$

$$\log \Re_{d} = n_{d} \log (1 - \Omega_{cal}) + \log k_{d}$$
$$\log \Re_{p} = n_{p} \log (\Omega_{cal} - 1) + \log k_{p}$$

• They can vary significantly

 $\Re_{d}(\% \text{ day}^{-1})_{\text{Indian ocean}} = 10^{4.3}(1 - \Omega_{\text{cal}})^{5.2}$

 $\Re_{d}(\% \text{ day}^{-1})_{\text{Pacific ocean}} = 10^{2.7}(1 - \Omega_{\text{cal}})^{3.0}$

 $\Re_d(\% \text{ day}^{-1})_{\text{Atlantic ocean}} = 10^{3.1}(1 - \Omega_{\text{cal}})^{4.5}$

• Effect of particle size or mineral surface condition?

Mechanical model

The dissolution mechanism is pH dependent



Fig. 9.11 Experimentally determined dissolution rates of calcite (Iceland spar) as a function of bulk fluid pH and $P_{\rm CO_2}$ at 25°C. The results can be divided into three regions: a pH-dependent regime (Region 1), a transition region (Region 2), and a pH-independent region (Region 3) (see text for explanation). The boundaries between the regimes are shown for dissolution at $P_{\rm CO_2} = 1$ atm but, in general, the boundaries are a function of $P_{\rm CO_2}$ and reaction progress. (After Plummer *et al.*, 1978, Figure 1, p. 186.)



Processes on the surface of calcite

- 1. $CaCO_3 + H^+ \xrightarrow{k_1} Ca^{2+} + HCO_3^-$
- 2. $CaCO_3 + H_2CO_3^* \xrightarrow{k_2} Ca^{2+} + 2HCO_3^-$
- 3. $CaCO_3 + H_2O \xrightarrow{k_3} Ca^{2+} + CO_3^{2-} + H^+ \rightarrow Ca^{2+} + HCO_3^- + OH^-$
- 4. $Ca^{2+} + HCO_3^{-} \xrightarrow{k_4} CaCO_3 + H^+$



$$\Re = k_1 a_{H^+} + k_2 a_{H_2 CO_3^*} + k_3 a_{H_2 O} - k_4 a_{Ca^{2+}} a_{HCO_3^-}$$

Driven by diffusion of H+ to the surface pH < 5

Controlled by diffusion of products from the surface pH > 5

Extension to the natural environment

A RESERVOIR AS A REACTOR

Rezervoár



By Willi





Reservoir balance

$$A = \sum_{i} I + \sum_{i} P - \sum_{i} O - \sum_{i} L$$

A ... increase/decrease of substance in the system

- I... input
- P... production
- *O* ... output
- L... loss/consumption



Balance

• If A is zero, the system is in dynamic equilibrium (steady state)

- Will the content of the system grow?
- Will system content go down?
- How long will it take to reach balance?

Retention time in the reactor

- The average time spent by the particle in the system
- At a steady state, transport by advection only

$$\tau = \frac{\text{celková hmota}}{\text{hmotová výměna}} = \frac{M_{tot}}{dM_{in(out)}/dt} = \frac{V}{F}$$

- au ... retention time
- M_{tot} ... total content of the reactor
- $M_{in(out)}$... quantity entering/exiting the reactor
- *V*... reactor volume
- *F*... flux to/from the reactor (they are in equilibrium)



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- RYAN, Peter Crowley. Environmental and lowtemperature geochemistry. Chichester, West Sussex, UK: Wiley Blackwell, 2014. ISBN 978-1-118-86735-8.